

Cenomanian algae and microencrusters from the El Abra Formation, W Valles–San Luis Potosí, Mexico

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Abstract

This work deals with calcareous algae and microencrusters (*incertae sedis*) from the upper part of the El Abra Formation, which crops out at the W Valles–San Luis Potosí Platform. Two assemblages of calcareous algae are recognized: one composed of rivulariacean-type cyanobacteria such as *Cayeuxia kurdistanensis* Elliott and *Garwoodia* sp. as well as the dasycladalean species *Clypeina* sp., *Neomeris cretacea* Steinmann, *Salpingoporella hasi* Conrad, Radoičić and Rey, *Trinocladus* cf. *T. divnae* Radoičić; the bryopsidales *Boueina pygmaea* Pia, *Halimeda* cf. *H. ellioti* Conard and Rioult, and the microencrusters *Lithocodium aggregatum* Elliot and *Thaumatoporella parvovesiculifera* (Raineri), and the other assemblage consisting of the gymnocodacean algae *Permocalculus irenae* Elliott and *P. budaensis* Johnson. The first assemblage is associated with the benthic foraminifera *Pseudolituonella reicheli* Marie, *Hemicyclammina sigali* Maync, *Daxia cenomana* Cuvillier and Szakall, *Merlingina cretacea* Hamoui and Saint Marc, is dated as mid-late Cenomanian on the basis of the stratigraphic distribution of the foraminifera, and suggests an oligotrophic environment within the euphotic zone on the open marine platform. Over this is documented a succession consisting of a community with abundant gymnocodacean algae and a decrease in the benthic foraminifera which could be related to increased nutrient availability in mesotrophic conditions. The shallow-water association composed of the studied microfossils is typical of the Tethys Realm.

Keywords: Algae, microencrusters, El Abra Formation, Cenomanian, Valles–San Luis Potosí Platform, Mexico.

Resumen

Este trabajo trata sobre las algas calcáreas y microencrustantes (*incertae sedis*) de la parte superior de la Formación El Abra que aflora en la parte occidental de la Plataforma Valles-San Luis Potosí. Se reconocen dos conjuntos de algas calcáreas, uno compuesto por cianobacterias tales como *Cayeuxia kurdistanensis* Elliott, *Garwoodia* sp. así como dasycladaleans: *Clypeina* sp., *Neomeris cretacea* Steinmann, *Salpingoporella hasi* Conrad, Radoičić y Rey *Trinocladus* cf. *T. divnae* Radoičić, bryopsidales: *Boueina pygmaea* Pia, *Halimeda* cf. *H. ellioti* Conard y Rioult y los microencrustantes *Lithocodium aggregatum* Elliott *Thaumatoporella parvovesiculifera* (Raineri). La otra asociación consiste de algas gymnocodáceas como: *Permocalculus budaensis* Johnson y *P. irenae* Elliott. El primer grupo se asocia con los foraminíferos bentónicos *Pseudolituonella reicheli* Marie, *Hemicyclammina sigali* Maync, *Daxia cenomana* Cuvillier y Szakall, *Merlingina cretacea* Hamoui y Saint Marc, lo cual es datado como Cenomaniano medio-tardío sobre la base de la distribución estratigráfica de los foraminíferos. Este conjunto sugiere un ambiente oligotrópico dentro de la zona eufótica en la plataforma marina abierta. Sobre esta sucesión se documenta una comunidad con algas gymnocodáceas abundantes y una disminución en los foraminíferos bentónicos, lo cual podría estar relacionado con un aumento de la disponibilidad de nutrientes en condiciones mesotróficas. La asociación de aguas someras compuesta de los microfósiles estudiados es típica del dominio Tethysiano.

Palabras clave: Algas, microencrustantes, Formación El Abra, Cenomaniano, Plataforma Valles-San Luis de Potosí, México.

1. Introduction

The Valles-San Luis Potosí Platform (VSLPP) is part of a large carbonate platform system that rimmed the ancestral Gulf of Mexico during the Mid-Upper Cretaceous (Scott, 1990). The shallow-water deposit is represented by the El Abra Formation, which holds a rich assemblage of benthic foraminifers. Together with benthic foraminifers the calcareous algae are an important element in the Cretaceous shallow-water deposits. The term “El Abra Limestone” was first used by Garfias (1915) for the shallow-water carbonates of the Sierra de El Abra. This unit is characterized by a great variety of invertebrate fossils, mostly Albian rudists, which have been analysed and published by authors including Coogan (1973), Alencaster (1987, 1998), and Alencaster and García Barrera (2008).

The benthic foraminifers from the El Abra Formation were reported by Bonet (1956); Tavitas and Solano (1984); Omaña and Torres Hernández (2000); Ornelas *et al.* (2006). In Mexico, few studies of fossil algae have been carried out; *Neomeris cretacea* was described for the first time by Steinmann (1899) from the Cenomanian of the Cerro Escamela in the state of Veracruz. Barattolo (1983) described *Triploporella steinmanni* (probably of mid-Albian age) of the Cretaceous of Orizaba from the same locality of Steinmann (1899); later, Barattolo (1990) made a careful revision of *Neomeris cretacea*.

Other reports on Cretaceous algae have been published by Buitrón *et al.* (1995) about the Albian–Cenomanian algae from the El Abra Limestone. Hernández-Romano *et al.* (1998); Aguilera Franco *et al.* (2001) recorded algae in the Cenomanian from the Guerrero–Morelos Platform. In the Chiapas region, Michaud (1987) and Deloffre *et al.* (1985) registered the occurrence of Late Cretaceous algae. Filkorn and Scott (2011) found a late Albian algal assemblage from the Mal Paso Formation in the state of Guerrero.

The purpose of this study is to document the occurrence of rivulariacean-type cyanobacteria, calcareous algae (dasycladales, bryopsidales) and microencrusters from the upper part of the El Abra Formation in material from the west part of the Valles–San Luis Potosí Platform, and to describe and illustrate the species identified.

2. Geological Setting

The study area is located east of the city of San Luis Potosí (Figure 1). This area is situated on the western part of the Valles–San Luis Potosí Platform (VSLPP). The Early Cretaceous was a time of remarkable tectonic stability in the Gulf of Mexico Basin, characterized by decreased terrigenous influx and the development of stable shelves, ramps and platforms bordering the deep central part of the Gulf of Mexico basin, which became the site of widespread carbonate deposition, particularly during the Albian. Along the western flank of the basin in east-central

Mexico, carbonate platforms were restricted to more local developments such as the VSLPP, which remained active until the earliest Late Cretaceous (Cenomanian) (Salvador, 1991).

3. Material and Methods

The material studied comes from the upper part of the El Abra Limestone. It was collected from four localities situated at the western part of the VSLPP: Llano del Carmen (LLC), Guadalcázar (G), Cerritos (C), and El Temazcal (ET). The limestone was examined in thin sections and algae and microencrusters were studied. The micropaleontological preservation is good, permitting identification. Most of the specimens are identified and illustrated from cross-sections.

The biostratigraphical framework is based on the benthic foraminiferal stratigraphic ranges. In addition, a microfacies study was carried out to infer the paleoenvironment.

4. Results

4.1. Lithology and microfacies

The samples with algae were collected east of the city of San Luis Potosí and proceed from the El Abra Formation. This unit presents two facies: the Taninul and El Abra, which are stratigraphically correlated (Aguayo, 1998). In this study from the upper part of the El Abra Formation (Taninul Facies), two main microfacies were distinguished.

Microfacies 1 is a peloidal bioclastic packstone and well sorted grainstone (Figure 2a-b).

This microfacies includes an algal assemblage composed of *Chypeina* sp., *Neomeris cretacea* Steinmann, *Salpingoporella hasi* Conrad, Radoičić and Rey, *Trinocladus* cf. *T. divnae* Radoičić, *Boueina pygmaea* Pia, and *Halimeda* cf. *H. elliotti* Conard and Rioult, *Cayeuxia kurdistanensis* Elliott, *Garwoodia* sp. and the microencrusters *Lithocodium aggregatum* Elliott and *Thaumatoporella parvovesiculifera* (Raineri). In this interval we identified the following benthic foraminifera: *Pseudolituonella reicheli*, *Hemicyclammia sigali*, *Daxia cenomana*, *Merlingina cretacea*, *Cuneolina conica*, *Cuneolina parva*, *Pseudocyclammia* sp., *Nezzazata simplex*, *Peneroplis parvus*, *Dicyclina schlumbergeri*, *Minouxia inflata*, *Nezzatinella picardi*, *Spiroloculina cretacea*.

Other fossil components of this microfacies include fragments of rudists, gastropods, corals, echinoderms, and calcareous worm tubes. This microfacies was recorded at all four localities (LLC, G, C, and ET). The abundant and diversified fauna and flora and the packstone texture of the rock which characterizes microfacies 1 suggest a depositional environment within the euphotic zone on the shallow open marine platform (zone 7 Flügel, 2004; Wilson, 1975). The environmental interpretation is also supported by

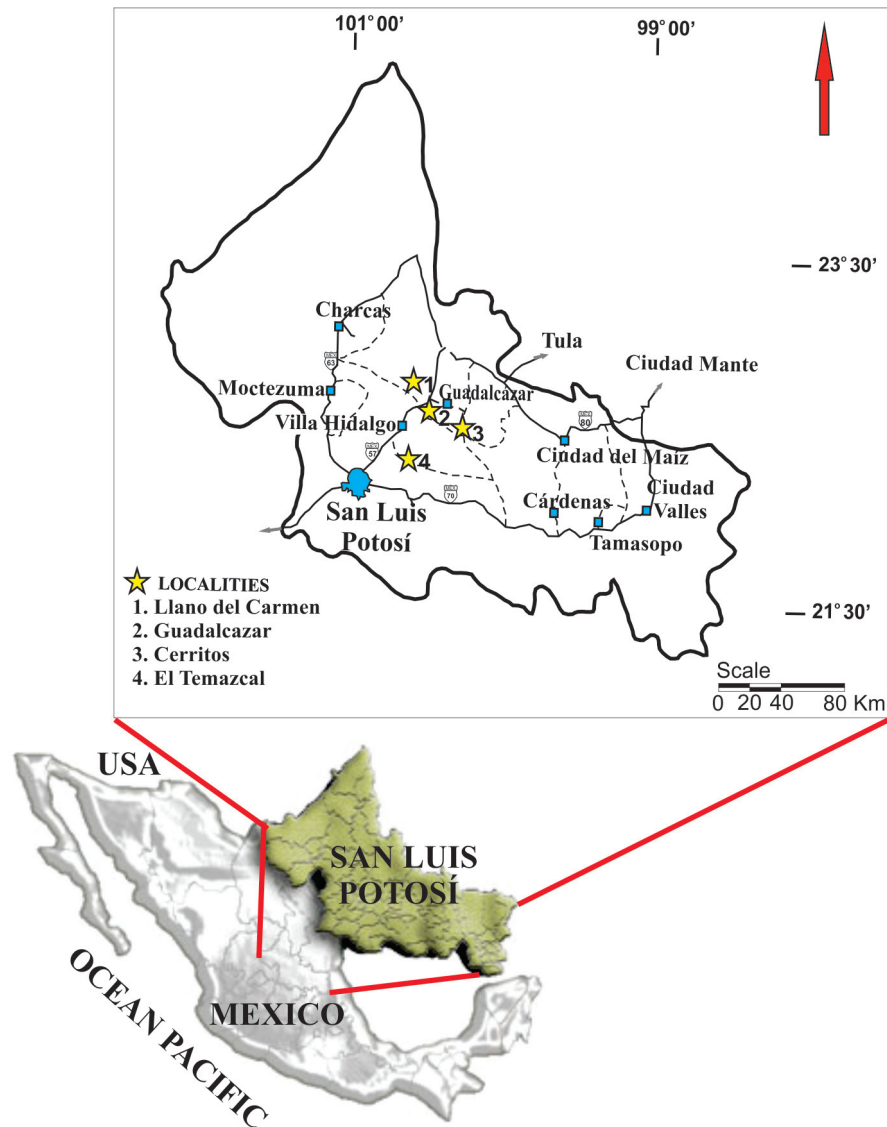


Figure 1. Geographic map with the location of the studied area.

the presence of calcareous algae and benthic foraminifera.

Microfacies 2 is an algal-foraminiferal wackestone-packstone characterized by fragments of *Permocalculus* (*Permocalculus irenae* Elliott, *P. budaensis* Johnson), and scarce benthic foraminifera such as *Nezzazatinella picardi*, miliolids, rotalids, textularids, gastropods, and echinoderms. This microfacies was identified in the Cerritos section (Figure 2c-d). The abundant occurrence of gymnocodaceans, as well as less diverse benthic foraminifera, indicates a marine deposit associated with an increase in the flux of nutrients (mesotrophic stage) before platform flooding.

4.2. Age

The benthic foraminiferal association is used for defining the age in the succession studied. On the basis of the

stratigraphic distribution of *Pseudolituonella reicheli*, *Daxia cenomana*, *Merlingina cretacea* and *Hemicyclammina sigali*, the interval that contains the algae and microencrusters was dated as mid-late Cenomanian age. This dating is reinforced by the occurrence of the *Whiteinella cretacea* Zone (latest Cenomanian–earliest Turonian) that overlies the upper part of the El Abra Formation (Omaña *et al.*, 2013).

4.3. Systematic Paleontology

The thin sections that contain the algae are housed in the Paleontology Collection of the Instituto de Geología (Universidad Nacional Autónoma de México).

PHYLUM CYANOPHYTA Sachs, 1874
Order Nostocales Geitler, 1925
Family Rivulariaceae Robenhorst, 1865

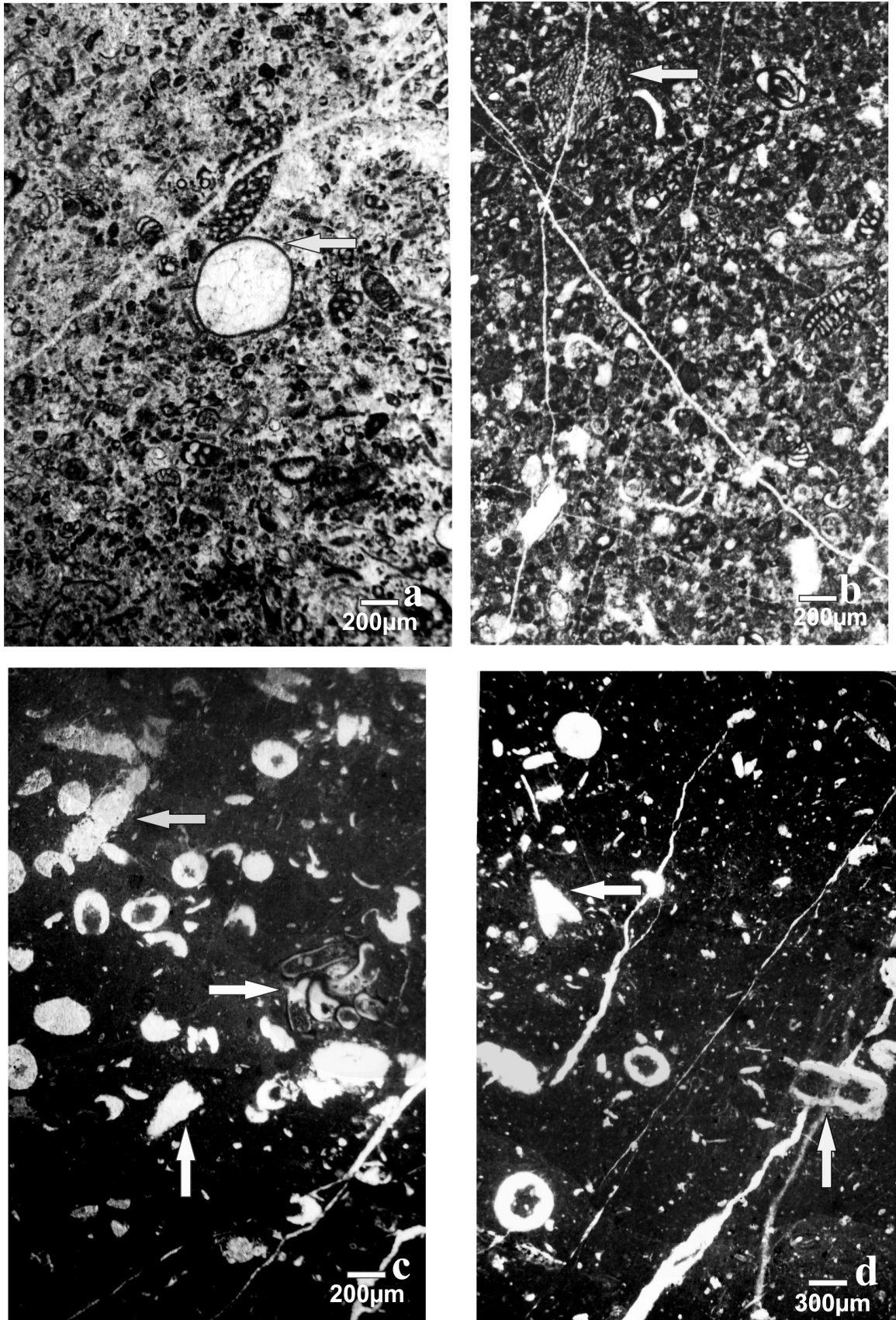


Figure 2. Microfacies with algae, foraminifers and encrusting associations. a) Foraminiferal packstone with *Thaumatoporella vesicularis* (Sample C-4). b) Peloidal bioclastic packstone with algae and foraminifers (Sample). c) Gymnocodacean algal packstone-wackestone with serpulids (Sample C-8). d) Gymnocodacean algal wackestone (Sample C-8).

The Cyanobacteria are an especially difficult group to classify. They are possibly one of the most ancient groups of organisms on earth, with some fossil representatives having very similar morphologies to present-day species (Schopf, 1974; Knoll, 2008).

Genus *Cayeuxia*

Frollo (1938) proposed the genus *Cayeuxia* with two new species, *C. piae* and *C. moldavica*, regarded as Codiaceae from the Tithonian, in East Carpathians. Dragastan (1985) transferred the genus to the cyanophyceans as a synonym of the genus *Rivularia*.

Dragastan (2008) noted that Luchinina and Shuysky in Chiuvashev *et al.* (1987) introduced other classification Phylum Cyanophyta, Family Garwoodiaceae including *Ortonella*, *Garwoodia*, *Hedstroemia*, *Cayeuxia*.

Type species. *Cayeuxia piae* Frollo, 1938.

Cayeuxia kurdistanensis Elliott, 1957

Figure 3 a, b, c

1957 *Cayeuxia kurdistanensis* Elliott, p. 790–791, pl. 25, figs. 8–10; Johnson, 1969, 37–38, pl. 26, fig. 1; Buitrón *et al.*, 1995, p. 150, pl. 3, fig. 1; Filkorn and Scott, 2010, p. 182, fig. 3, 3; *Rivularia kurdistanensis* (Elliott), Dragastan, 1985, pl. 4, figs. 8–14; Mancinelli and Ferrandes, 2001, p. 538, figs. 3.4, 4.6.

Description. Fan-like thallus formed by radiating tubular filaments that diverge toward the distal margin and bifurcate from preceding tubes with an angle of divergence of about 45°, which is characteristic of *Cayeuxia*; the tubes filled with spar calcite as Filkorn and Scott (2011) have already observed.

Stratigraphic distribution. *Cayeuxia kurdistanensis* was described for the first time by Elliott (1957) from the Aptian of Iraq. This species has a wide distribution in the Jurassic and Cretaceous strata in Europe (Italy and France) as well as in the Americas (USA, Mexico, and Argentina).

Measurements.

Height 1mm

Occurrence. *Cayeuxia kurdistanensis* identified from the Llanos del Carmen (Sample LLC-1).

Garwoodia sp. Figure 3 d

Description. Thallus with tubes dichotomically branched laterally at an angle of about 90 degrees; the sheaths arranged in an irregular form; their walls are separated and the space is filled with calcite.

Remarks. The coarser tubes of *Garwoodia* have been used as an argument for assigning it to the udotacean green algae (Flügel, 2004).

Measurements.

Length 1000 µm.

Width 700 µm.

PHYLUM CHLOROPHYTA Pascher, 1914, p. 147

Class Dasycladophyceae Hoek *et al.*, 1995

Order Dasycladales Pascher, 1931

Family Polyphysaceae Kützing, 1841

Genus *Clypeina* Michelin, 1845

According to Granier *et al.* (2014) the type species of *Clypeina* is *Clypeina marginiporella* (Michelin, 1845) which was originally described as an organism similar to a coral by Michelin (1840–1847; p. 177–178). The genus name derives from the Latin word “*clypeus*” for the reason that the first specimens collected by Michelin had the appearance of “perfect rings” and “half rings.” In the second half of the nineteenth century, there was no consensus on the taxonomic position of these isolated “whorls” which were placed into the foraminifers by various authors as Parker and Jones (1860, p. 473–474); Carpenter (1862, p. 130–131), and Gümbel (1872, p. 262), but they were appropriately recognized as algae by Munier-Chalmas (1877) who regarded them as “*Siphonées verticillées*”. Morellet and Morellet (1913) put them in the “*Acétabulariées*”. In 1918, these authors proposed a new species: *Clypeina helvetica*, in addition to the two species previously identified from the Cenozoic, *C. marginiporella* (Michelin, 1845) and *C. digitata* (Parker and Jones, 1860). However, there was no record of Mesozoic forms up to the description of *Clypeina jurassica* from the Purbeckian strata (Favre and Richard, 1927).

Clypeina sp.

Figure 4 a,b,c,d, e, f, g

Description. Tallus cylindrical, euspondyl composed of one fertile whorl that consists of seven to 13 elongated primary laterals irregularly arranged containing one ovoid sporangial structure. The laterals are connected to the stem by small pores. The outer thallus diameter (D) was varying from 1000 to 625 µm; inner thallus diameter (d) 450–250 µm; pore diameter (p) 250–50 µm.

Class Dasycladophyceae Hoek *et al.*, 1995

Order Dasycladales Pascher, 1931

Family Dasycladaceae Kützing, 1841

Tribe Dasycladeae Pia, 1920

Genus *Neomeris* Lamouroux, 1816

According to Granier *et al.* (2012a) “The main feature of the *Neomeris* genus is a main stalk bearing regularly spaced

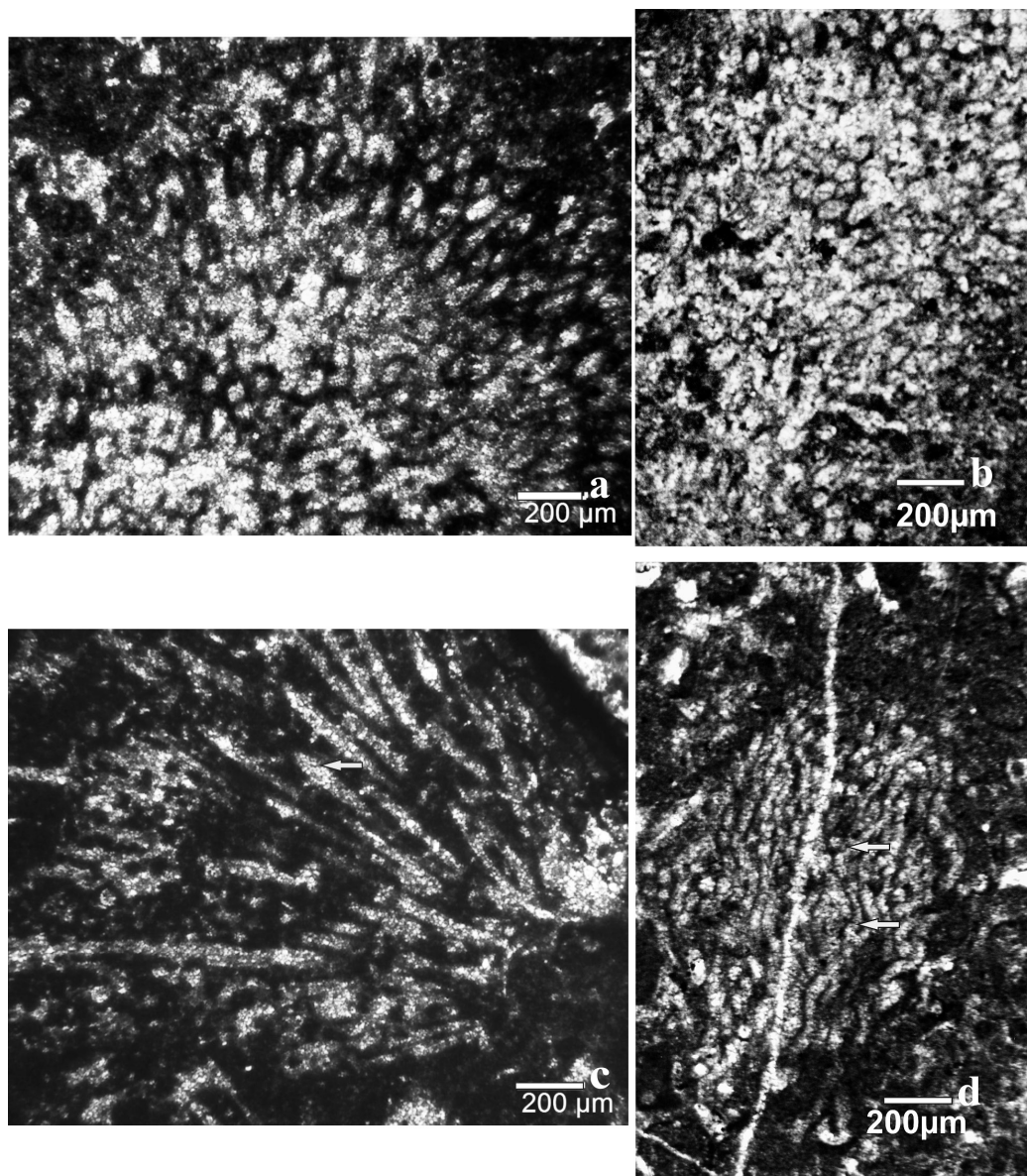


Figure 3. Cenomanian rivulariacean-type cyanobacteria from the El Abra Formation. a) Longitudinal section of *Cayeuxia kurdistanensis* (Sample LLC-1). b–c) Tangential section of *Cayeuxia kurdistanensis* (Sample LLC-1). d) Longitudinal section of *Garwoodia* sp. (Sample LLC-2).

verticils of laterals with two (exceptionally three, e.g., in the living *N. stipitata* Howe, 1909) secondary sterile segments and, where they are fertile, one gametophore in terminal position (choristosporate type) per primary segment”.

Neomeris cretacea Steinmann, 1899

Figure 5 a,b

Neomeris (Herouvalina) cretacea Steinmann, 1899, p. 149–154, figs. 14–20; Raineri 1922, p. 74–75, figs. 5, 6; Basson and Edgell, 1971, p. 422, pl. 5, figs. 4, 5; Kuss, 1986 b, p. 228–230, fig. 5d; Kuss and Conrad, 1991, p. 872, fig. 2.17.

Description. Cylindrical, elongated thallus with a

central hollow which have two types of ramifications: primary branches each bearing two sterile secondary branches which are situate on both sides of the fertile ovoid to ellipsoidal ampulla.

Measurements.

Outer thallus diameter (D) 650 µm.

Inner thallus diameter (d) 350 µm.

Sporangia diameter (p) 50 µm.

Remarks. *Neomeris cretacea* was first described by Steinmann (1899) from the Cenomanian of Cerro Escamela in Mexico.

Stratigraphic distribution. The earliest representative of *Neomeris* should have been “born” at the transition between the Jurassic and the Cretaceous, considering the

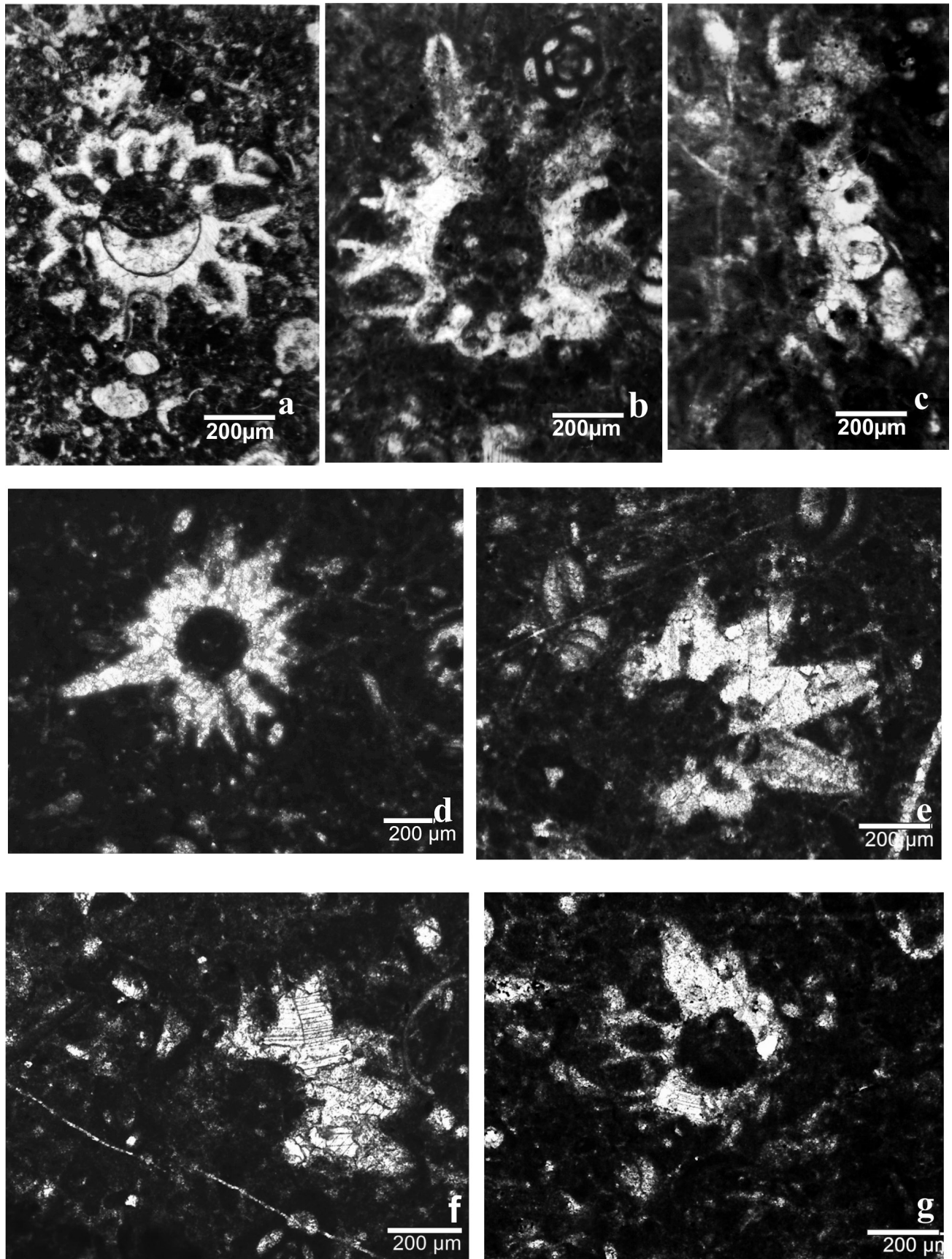


Figure 4. Cenomanian Dasycladacean algae from the El Abra Formation. a) Cross section of *Clypeina* sp. (Sample G-5). b, d, g) Cross section of *Clypeina* sp. (Sample G-5). c) Tangential section of a verticil of *Clypeina* sp. (Sample G-5). e–f) Cross section of *Clypeina* sp. (Sample G-5).

choristosporate model of reproduction, which is the main trait of the family Dasycladaceae, as emended by Granier and Bucur (Granier *et al.*, 2012b).

Other reports are from the Late Cretaceous (Cenomanian-Turonian) of Iraq. In Lebanon it spans from late Aptian to Cenomanian, but the most typical specimens recorded are from Cenomanian (Basson and Edgell, 1971).

Occurrence. In our material we found this species only in the Temazcal (Sample ET-4) associated with *Daxia cenomana*, *Cuneolina pavonia*, *Cuneolina parva* and *Dicyclina schlumbergeri*.

Family Triploporellaceae Pia, 1920

Tribe Salpingoporelleae Bassoullet *et al.*, 1979

Sub-tribe Salpingoporellinae Bassoullet *et al.*, 1979

Genus *Salpingoporella* Pia, 1918 in Trauth, emend Carras *et al.*, 2006

Type-species. *Salpingoporella muehlbergii* (Lorenz, 1902) Pia, in Trauth, 1918 emend. Carras *et al.*, 2006.

Salpingoporella hasi Conrad, Radoičić and Rey, 1977
Figure 5 c

Salpingoporella hasi Conrad, Radoičić and Rey, 1977, pl. 1, fig. 5–6; Cherchi and Schroeder, 1980, pl. 1, 474; Kuss and Conrad, 1991, fig. 4.17; Radoičić, 1994, pl. 1, fig. 1; Sokač, 1996, pl. 2, fig. 1–5; pl. 8; Masse and Arnaud-Vanneau, 1999, pl. 1, fig. 11–12; Masse and Isintek, 2000, pl. 2, fig. 3–5; Radoičić, 2006, p.85, pl. 7, fig. 1

Original description. “Cylindrical thallus whose verticils are made up of a small number of branches, transversally flattened, and whose tubular proximal part is well developed. At their distal end, the branches widen out quickly and become rectangular. The nearby branches belonging to the same verticil may touch each other or not, depending on the space available, which in turn depends on their length and on the diameter of the siphon. The calcareous envelope is made up of a simple mosaic of hyaline crystals” (Carras *et al.*, 2006).

Measurements.

Outer thallus diameter (D) 450 µm.

Inner thallus diameter (d) 250 µm.

Stratigraphic distribution. *Salpingoporella hasi* was described as an Albian-Cenomanian species of different localities of the Mediterranean region (Conrad *et al.*, 1977). This species has been recorded in the Albian from the Karaburun Peninsula in Turkey (Masse and Isintek, 2000) and the Guyot Resolution in W Pacific (Masse and Arnaud-Vanneau, 1999).

Occurrence. *Salpingoporella hasi* identified in the Guadalcazar site (Samples G-5).

Family Triploporellaceae (Pia, 1920) emend. Berger and Kaeffer, 1992

Subtribe Triploporellinae (Pia, 1920) emend. Bassoullet *et al.*, 1979

Genus *Trinocladus* Raineri, 1922

Elliott (1972, p. 619) indicated that *Trinocladus* is composed of “successive verticils of radial branches, each branch showing outwardly widening primaries giving rise to several secondaries, and these in turn to bunches of tertiaries. Branches of the lower verticils may not show the full detail. Branches usually not alternate in position from verticil to verticil.”

Trinocladus cf. *T. divinae* Radoičić, 2006
Figure 5 e

Description. Fragment of longitudinal section with typical *Trinocladus* organization of the laterals with inflated parts of the secondaries that are well preserved and the occurrence of small tertiaries.

Measurements.

Length of segment 1250 µm.

Occurrence. *Trinocladus* cf. *T. divinae* identified in the Guadalcazar site (Samples G-5).

Bryopsidales

Udotaceae

The Udotaceae in the Cretaceous are represented by assemblages of dissociated segments which have been referred to three genera, *Boueina*, *Arabicodium*, and Halimedaceae *Halimeda* (Bassoullet *et al.*, 1983). Although the segments are different in size and shape and internal structure, they all coincide in possessing a longitudinal medullary zone and a cortical zone which is similar to that in living *Halimeda* and closely related to it.

Steinmann (1901) stated that *Boueina* can also be differentiated from *Halimeda* based on its external character. He claimed “that *Boueina* is a cylindrical, unbranching plant, then we may emphasize that this feature is not common for the genus *Halimeda*, in spite of the similarities concerning the internal structure. All *Halimeda* species are branched and articulated.”

Elliott (1965, p. 199) stated that *Boueina* “in thin-section shows a medullar zone of coarse tangled threads, and a cortex of approximately radial finer branching threads, whereas in *Halimeda* the central coarse threads are mostly longitudinally directed and the cortical threads show constrictions and swellings.” This last characteristic underlined by Steinmann (1901), Elliott (1965), and Bucur (1994) has been omitted by some subsequent authors who studied algae from the *Boueina-Arabicodium-Halimeda* group.

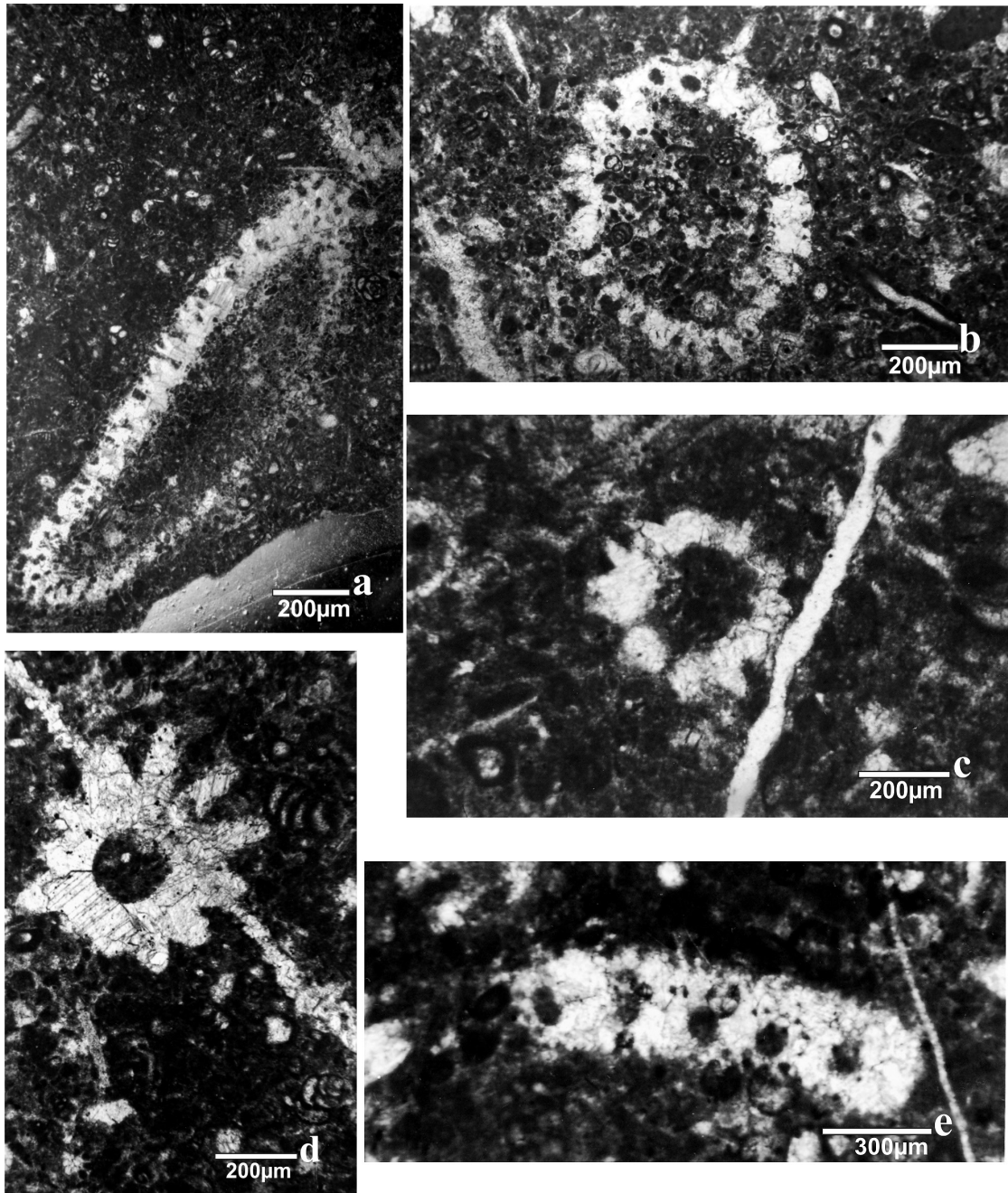


Figure 5- Cenomanian Dasycladacean algae from the El Abra Formation. a) Tangential section of *Neomeris cretacea* (Sample ET-4). b) Cross section of *Neomeris cretacea* (Sample ET-4). c) Cross section of *Salpingoporella hasi* (Sample G-5). d) Cross section of *Salpingoporella* sp. (Sample G-5). e) Longitudinal section of *Trinocladus* cf. *T. divnae* (Sample G-5).

Dragastan and Herbig (2007) indicated that “the classification of fossil-calcified-green siphonaceous algae is based on skeleton morphology. In general, internal morphology shows siphons, arranged in a central medulla and external cortex. Descriptive terms have been used with different meanings, such as tubes or filaments.” Littler and Littler (1990, 1992) point out that the filament term, used by previous workers, is generally defined as a chain of cells and is incorrectly applied to siphonaceous algae such as *Udotea*, *Halimeda*, and *Penicillus*. They proposed the siphon terms

for the medulla and utricle for the cortex.

Class Bryopsidophyceae Round, 1963

Order Bryopsidales Schaffner, 1922

Suborder Halimedineae Hillis-Colinvaux, 1984

Family Udoteaceae (Endlicher) Agardh, 1887–1888

Genus *Boueina* Toula, 1884

Type-species. *B. hochstetteri* Toula, 1884

Boueina pygmaea Pia, 1936

Figure 6 a

Boueina pygmaea Pia, 1936, p. 12–13, pl. 5, figs. 1–9; Elliot, 1965, p. 204, pl. 24, figs. 2, 6; Kuss, 1986, p. 231, fig. 5f-i; Kuss, 1994, p. 313, fig. 9; Bauer et al., 2002, p. 17, pl. 4, fig. 3.

Description. Cylindrical thalli showing a strong calcification of the medullary and cortical zones. The interwoven siphons of the narrow medullary zone diverge to the cortical zone; toward the periphery where they then ramify into smaller utricles curving and diverging in all directions.

Measurements.

Height of thalli (h) 1250 µm.

Width of thalli 750 µm.

Medular zone 300 µm; cortical zone 450 µm.

Stratigraphic distribution. Pia (1936) described *Boueina pygmaea* from the Cenomanian-Turonian of Lybia. It has been recorded from the Cenomanian of France and Spain (Pfender, 1940; Ramírez del Pozo and López, 1988); Kuss (1994) found the species from the Cenomanian limestone of Gebel El Musherah in northern Sinai. Elliot (1965) reported *Boueina* cf. *B. pygmaea* from the Cenomanian of Bou-Saâda, Algeria.

Occurrence. We found *Boueina pygmaea* in the Llanos del Carmen site (Sample LLC-1).

Suborder Halimedineae Hillis-Colinvaux, 1984

Family Halimedaceae Link, 1832

Genus *Halimeda* Lamouroux, 1812*Halimeda* cf. *H. ellioti* Conard and Rioult, 1977

Figure 6 b

Description. Cylindrical thalli, with elongated segments; large and thicker medullary zone crossed by tubular siphons that extend as bifurcated cortical utricles which are inflated and constrained in their bases (Bucur et al., 2010).

Remarks. *Halimeda* cf. *H. ellioti* identified in our material is identical to Figure 9 (plate 6) illustrated by Radoičić (2006).

Measurements.

Length of segment 1950 µm.

Diameter of the medullary zone 400 µm.

Diameter of the cortical zone 200 µm.

Stratigraphic distribution. *H. ellioti* was described by Conard and Rioult (1977) from the Upper Cretaceous (Turonian) of southern France. Some Cenomanian-Turonian series in the Pastrok Mountain area (Kukes Cretaceous Unit) in Albania are characterized by the presence of the little-known udoteacean species *Halimeda ellioti* Conard and Rioult. “The inferences based on the data presented are that the range of *Halimeda ellioti* is from the middle Cenomanian into the middle Turonian and its maximum development is in the uppermost Cenomanian and lower

Turonian” (Radoičić, 2005).

Occurrence. *Halimeda* cf. *H. ellioti* identified in the Llanos del Carmen locality (Sample LLC-1).

Family Gymnocodiaceae Elliot, 1955

Genus *Permocalculus* Elliot, 1955**Type-species.** *Permocalculus irenae* Elliot, 1955*Permocalculus budaensis* Johnson, 1968

Figure 6 c, e

Permocalculus budaensis Johnson, 1968, p. 8, 9, pl. 1, figs. 2–5; Hernández Romano et al., 1998, p. 50, fig. 5d.

Description. Segmented thallus with cortical crenulations, with the sporangia in cortical or subcortical position. The weak calcification is responsible for the darker appearance of *P. budaensis* in comparison with other species of *Permocalculus* (Kuss and Conrad, 1991).

Measurements.

Height of thalli (h) 1350 µm.

Diameter of thalli 600 µm.

Stratigraphic distribution. *Permocalculus budaensis* was described and illustrated from the Buda Limestone of Cenomanian age (Johnson, 1968). Kuss and Conrad (1991) recorded the occurrence of the species in the Cenomanian of Jordan and within the Turonian of the Sinai.

Occurrence. This species was recorded from the Cerritos locality (Sample C-8).

Permocalculus irenae Elliot, 1955

Figure 6 d

Permocalculus irenae Elliot, 1955, p. 258–259, pl. 47, fig. 2, pl. 48, figs. 2–6; Johnson, 1965, p. 719, pl. 89, fig. 4; Basson and Edgell, 1971, p. 429, pl. 6, fig. 8; Kuss and Schlaginweit, 1988, p. 92, pl. 19, fig. 1; Shirazi, 2008, p. 803, pl. III, figs. 1–9.

Description. Slightly segmented thallus, segments long, ovoid or subcylindrical, irregular finger-like or pinched and swollen units with fine pores at outer zone, internal sporangia are not visible. Calcification varies from thin to thick layers to total calcification. Broken thalli fragments are often accumulated, creating the typical “algal-debris-facies” of Elliott (1958).

Measurements.

Height of thalli (h) 450 µm.

Diameter of thalli 300 µm.

Stratigraphic distribution. Elliott (1958) described *Permocalculus irenae* from the Cenomanian of northern Iraq; Cenomanian of Iran and Albian-Cenomanian of Lebanon reported by Basson and Edgell (1971). Kuss and Conrad (1991) recorded the species in the Albian from the Sinai and the Cenomanian from Jordan and northeastern Egypt.

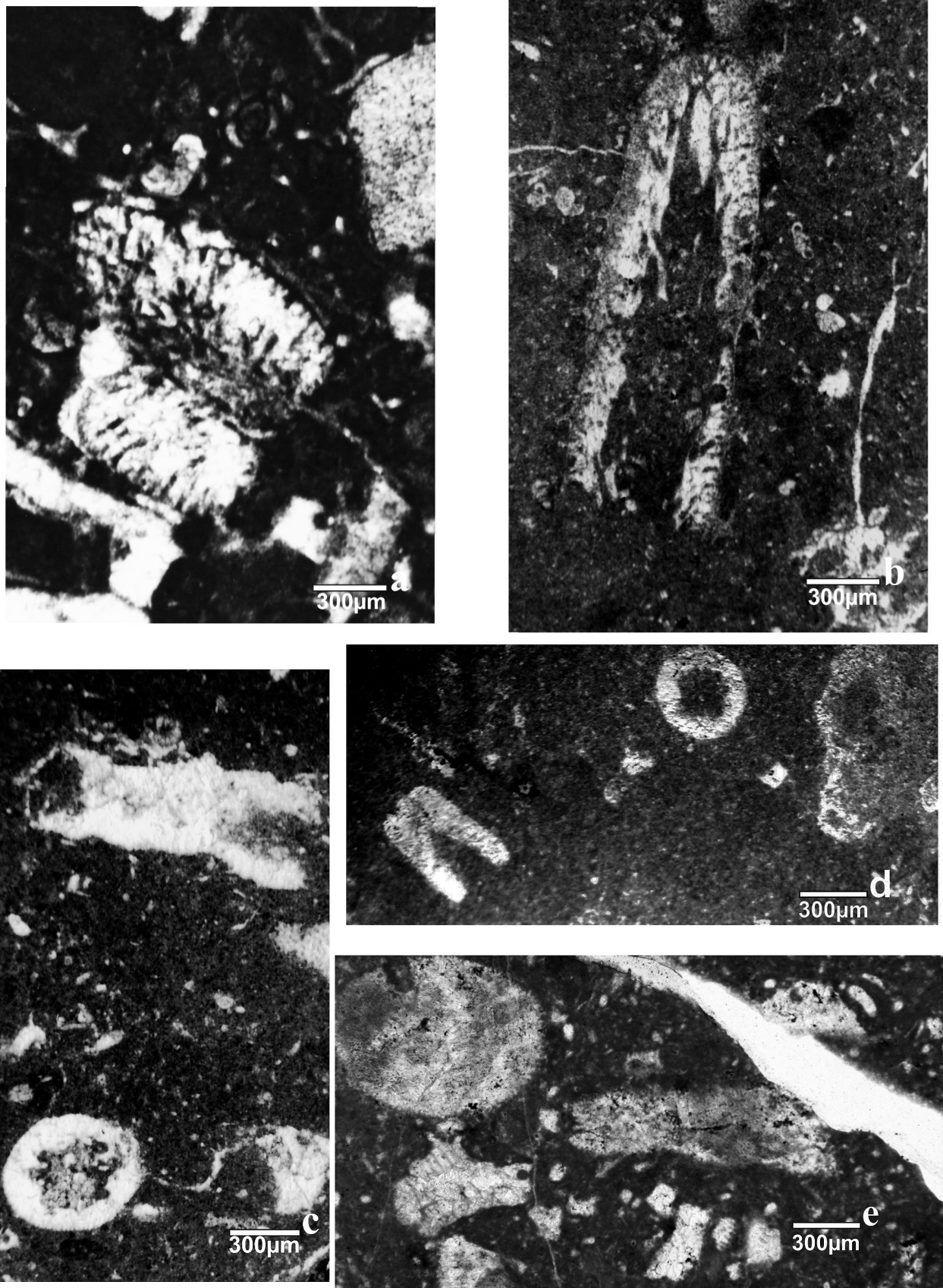


Figure 6. Cenomanian Bryopsidales and Gymnocodacean algae from the El Abra Formation. a) Longitudinal section of *Boueina pygmaea* (Sample LLC-4). b) Longitudinal section of *Halimeda* cf. *H. elliotti* (Sample LLC-1). c) Longitudinal and cross sections of *Permocalculus budaensis* (Sample C-8). d) Longitudinal and cross sections of *Permocalculus irenae* (Sample C-8). e) Longitudinal and cross sections of *Permocalculus budaensis* (Sample C-8)..

Occurrence. This species was recorded in Cerritos locality (Sample C-8).

Incertae sedis (Microencrusters)
Genus *Lithocodium* Elliot, 1956

Lithocodium aggregatum Elliot, 1956
Figure 7a, b

1956 *Lithocodium aggregatum* Elliot, p. 331, pl. 1, figs. 2, 4–5; Banner *et al.*, 1990, p. 24, pl. 1, fig. 2, p. 31, pl. 4 fig. 1; Radoičić, 2005, pl. 5, figs. 1–6; pl. 6, figs. 1, 2; Schlagintweit and Bover-Arnal, 2011, p. 4, figs. 2e, g.

Description. The *Lithocodium* crusts consist of an irregular network and the periphery presents a thin alveolar system forming bifurcated filaments perpendicular to the surface.

Remarks. *Lithocodium aggregatum* was described by Elliott (1956) from Iraq. The taxonomic position of this encrusting microorganism has been controversial during recent decades. *Lithocodium aggregatum* was originally described as a siphonal (=non-septate filaments) codiacean alga by Elliot (1956) and interpreted also as a codiacean green algae by Banner *et al.* (1990). Its codiacean nature was accepted by many subsequent authors (Praturlon, 1964; Johnson, 1964; Poignat, 1968; Radoičić, 2005). Besides the assumed green algal nature, other authors such as Koch *et al.* (2002) indicated that *Lithocodium aggregatum* is more similar to sponges.

Schlagintweit (2010) and Schlagintweit *et al.* (2010) considered *Lithocodium* to be a filamentous-septate heterotrichale ulvophycean alga (order Ulotrichales?) exhibiting a heteromorphic life cycle consisting of two phases: an epilithic gametophytic stage and an euendolithic sporophytic (*Gomontia*) stage.

The taxonomic interpretation of *Lithocodium aggregatum* as sessile green algae (Schlagintweit *et al.*, 2010) is accepted by Huck *et al.* (2012).

The taxon has been regarded as a foraminifer as well (Schmid and Leinfelder, 1995, 1996), which was discussed and rejected by Cherchi and Schroeder (2006) and Schlagintweit (2008); however, some authors, such as Dupraz and Strasser (2002), agree with the interpretation of Schmid and Leinfelder (1995, 1996) regarding the foraminiferous character of *Lithocodium aggregatum*. It has also been regarded as calcimicrobial colonies (Camoin, 1983). Cherchi and Schroeder (2006) showed that *Lithocodium aggregatum* is a colony of calcified cyanobacteria. This interpretation is confirmed by Conrad and Clavel (2008) who found that the *Lithocodium* crusts have a high organic content. Cherchi and Schroeder (2010, 2011, 2013) stated that *Lithocodium aggregatum* is a sponge (Entobia)-calcimicrobe consortium. They agree with the interpretation that these crusts are colonies of calcified microbes, probably cyanobacteria, which were

transformed into dense, homogeneous micrite. Numerous occasionally broken needle-shaped sponge spicules and minute sedimentary particles were introduced by currents, trapped by calcimicrobial mucilage, and incorporated into the crusts.

Stratigraphic distribution. *Lithocodium aggregatum* ranging to the Late Triassic–Middle Cretaceous carbonate platforms of the Tethyan realm (Elliott, 1963; Flügel, 2004)

Occurrence. Identified from the Llanos del Carmen (Sample LLC-1).

Genus *Thaumatoporella* Pia, 1927

Thaumatoporella parvovesiculifera (Raineri 1922)
Figure 7 d–e

1922 *Gyroporella parvovesiculifera* n. sp. Raineri, p. 83, pl. 13, fig. 17–18.

Thaumatoporella parvovesiculifera (Raineri, 1922): Pia, 1927, p. 69; Sartoni and Crescenti, 1962, p. 270, pl. 16; pl. 46, figs. 2–3, 5; De Castro, 2002, text-fig. 2, pl. 1, fig. 1–11; Flügel, 2004, p. 566, pl. 99, fig. 2; Senowbari-Daryan *et al.*, 2011, p. 512, fig. 11–M; Schlagintweit, 2013, p. 8, figs. 8 d–e.

Description. *Thaumatoporella parvovesiculifera* (Raineri, 1922) consists of an outer thin-walled structure with cells forming an irregular, lobed, ovoid to globular and cylindrical structure, with an internal cavity preserved as sparry calcite (Flügel, 2004).

Thaumatoporellaceans can be free-living, attached to hard substrates, or dwell as cryptoendoliths inside hollow bioclasts (Schlagintweit *et al.*, 2013).

Remarks. Raineri (1922) described *Gyroporella parvovesiculifera* from the Upper Cretaceous of southern Italy and placed it in the Dasycladacean algae. Later, Pia (1927) proposed the genus *Thaumatoporella* taking into account the differences between *Gyroporella* Raineri's genus, accepting its place within the Dasycladacean algae.

Pia (1938) observed that “the thalli were not growing in an upright position like that of the Dasycladacean but were creeping on the sea floor”. This author noted the peculiarity of this genus which has no internal structure and emphasized the high variability in morphology and dimension including the width of the thallus, pore diameter, and wall thickness.

This problematic genus has been interpreted as red algae (Ramalho, 1971; Flügel, 1979), chlorophycean algae (Barattolo, 1991) and as a different group of green algae (De Castro, 1990). *Thaumatoporella parvovesiculifera* can be free-living or attached to hard substrates, as has been already observed by Schlagintweit and Velić (2012).

Stratigraphic distribution. *Thaumatoporella parvovesiculifera* successfully inhabited different shallow-water environments over a long period of more than 150 million years from the Middle Triassic (Lindian) (de Castro, 1990; Flügel, 2004; Schlagintweit, 2012) to the lower Eocene (Ilerdian) in the Limalok Guyot (Central

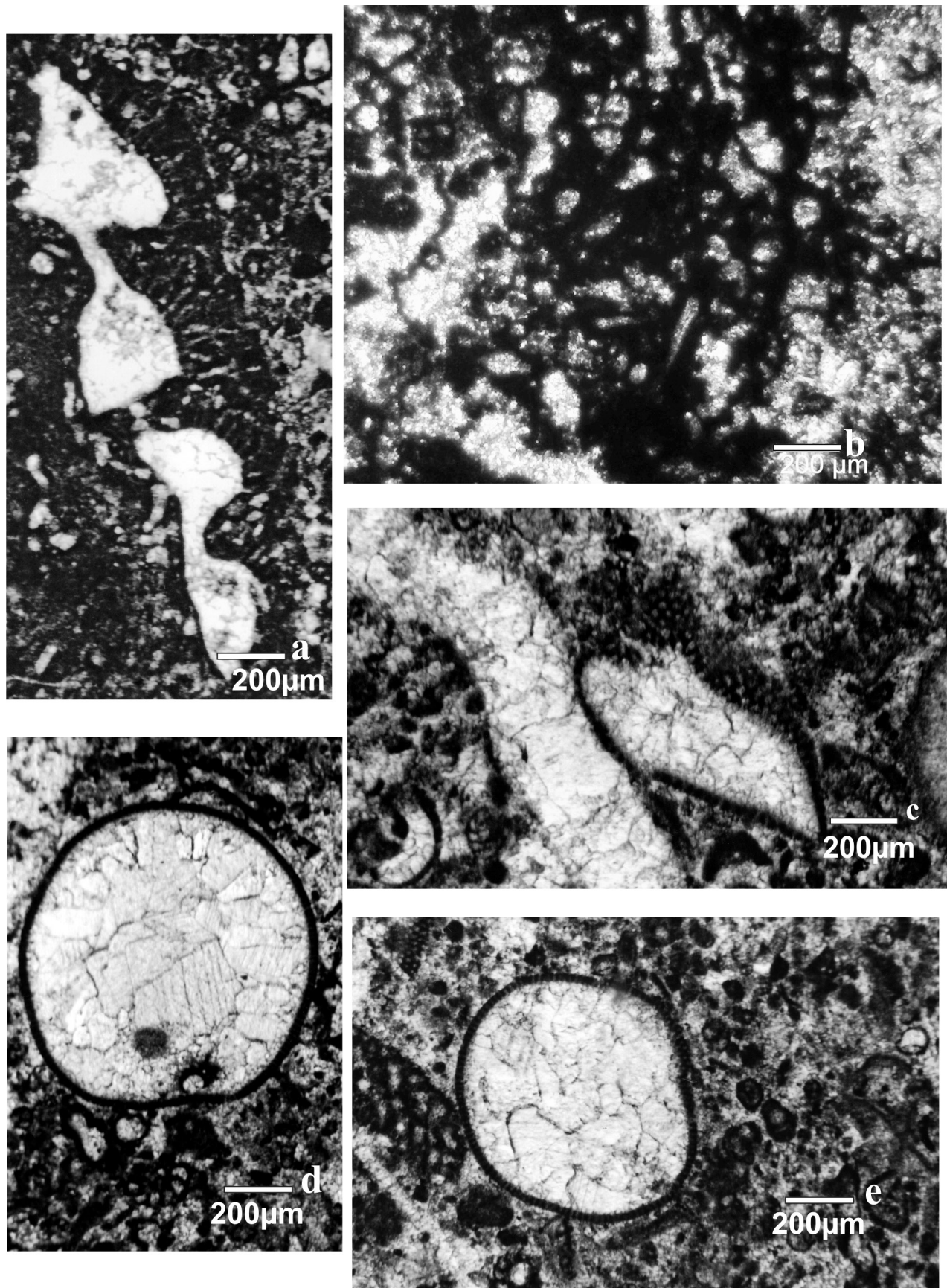


Figure 7. Cenomanian incertae sedis (microencrusters) from the El Abra Formation. a–b) *Lithocodium aggregatum* (Sample LLC-1). c–e) *Thaumatoporella parvovesiculifera* (Samples C-5, 6).

Pacific). At this site it was observed in younger sediments of middle Eocene (Lutetian) (Nicora *et al.*, 1995). Later, Vecchio and Barattolo (2006) and Vecchio and Hottinger (2007) recorded thamatoporellacean algae in southern Italy in the same interval; from these published data the occurrence of this fossil ranged to the Eocene (uppermost Ypresian–lowermost Lutetian).

Measurements.

Diameter 600–800 µm

Occurrence. *Thamatoporella parvovesiculifera* was living free or on a hard substrate, as was already observed by Schlagintweit and Velić (2012). In our material it is very common in the Cerritos locality (Samples C-5, C-6).

5. Conclusions

The deposit of the upper part of the El Abra Limestone (Taninul Facies) contains a diverse algal and benthic foraminiferal assemblage. It is dated as mid-late Cenomanian based on the occurrence of *Pseudolituonella reicheli*, *Daxia cenomana* and *Merlingina cretacea*.

The textural features and the microfossil association allow for two different depositional environments to be inferred. The first is characteristic of a warm shallow-water carbonate platform with open circulation and medium-high hydrodynamic energy with a great diversity of benthic and algal species in an oligotrophic stage, and the other represents a reduction of the benthic foraminiferal assemblages associated with the gymnocodacean algae such as *Permocalculus irenae* and *P. budaensis*, which could be related to increasing nutrient availability in mesotrophic conditions.

The algal association is typical Tethysian, similar to Mediterranean localities.

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